

# Explicit simulation of uncertainties in ocean models, and application in SANGOMA benchmarks

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**Since important decisions  
must rely on simulations,  
it is essential that its validity be tested,  
and that its advocates be able to describe  
the level of authentic representation  
which they achieved.**

Summer Computer Simulation Conference (1975),  
cited by Richard Hamming (1997)

# Outline

**1**

**Introduction**

**2**

**Explicit simulation of uncertainties**

**3**

**Stochastic circulation model**

**4**

**Stochastic ecosystem model**

**5**

**Stochastic sea ice model**

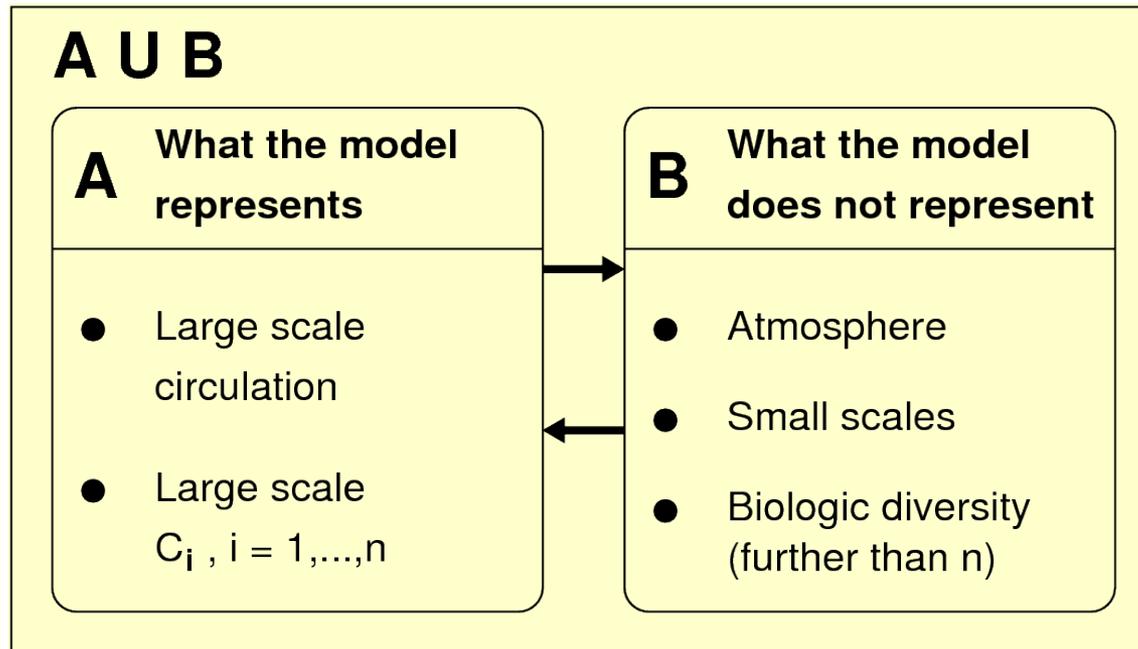
**6**

**Conclusions**

**1**

# **Introduction**

# Sources of uncertainties in ocean models



- Even if the dynamics of **U** can be assumed deterministic, the system **A** alone **cannot be assumed deterministic**.
- To obtain a deterministic model for **A**, one must assume, either that **B** is known ( $\rightarrow$  atmospheric forcing), or that the effect of **B** can be parameterized ( $\rightarrow$  paramétrisation of unresolved scales or unresolved biologic diversity).  
  
 $\rightarrow$  **B is the main source of uncertainty in the model.**

## Motivations for a probabilistic approach

**The deterministic approach is not always sufficient to describe the dynamical behaviour of the system**

**Comparison between simulations and observations is easier with the probabilistic approach**

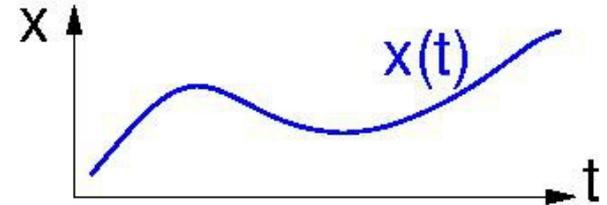
**A good knowledge of model accuracy is necessary to solve data assimilation problems**

# Probabilistic approach to ocean modeling

**Stochastic ocean dynamics**, explicitly simulating uncertainties

$$d\mathbf{x} = \mathcal{M}(\mathbf{x}, t)dt + \Sigma(\mathbf{x}, t)d\mathbf{W}_t$$

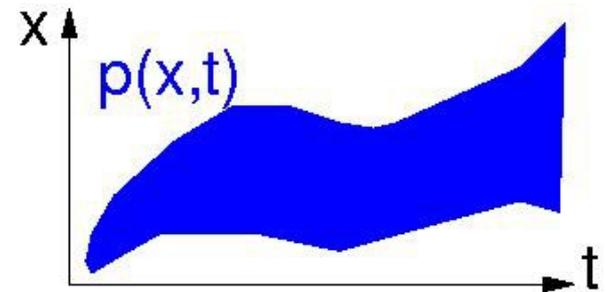
$$\text{où } \mathbf{x} = [x_1, \dots, x_N]$$



**Fokker-Planck equation**, for the probability distribution  $p(\mathbf{x}, t)$ ,  
following ideas at the origin of the Ensemble Kalman filter (Evensen, 1994)

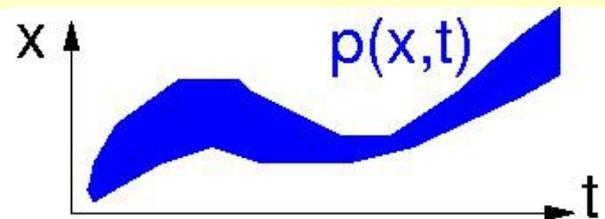
$$\frac{\partial p(\mathbf{x}, t)}{\partial t} = - \sum_{i=1}^N \frac{\partial}{\partial x_i} [\mathcal{M}_i(\mathbf{x}, t)p(\mathbf{x}, t)] \\ + \frac{1}{2} \sum_{i=1}^N \sum_{j=1}^N \frac{\partial^2}{\partial x_i \partial x_j} [D_{ij}(\mathbf{x}, t)p(\mathbf{x}, t)]$$

$$\text{où } \mathbf{D} = \Sigma\Sigma^T$$



**Conditioning to observations**, to reduce uncertainties

using an appropriate data  
assimilation method



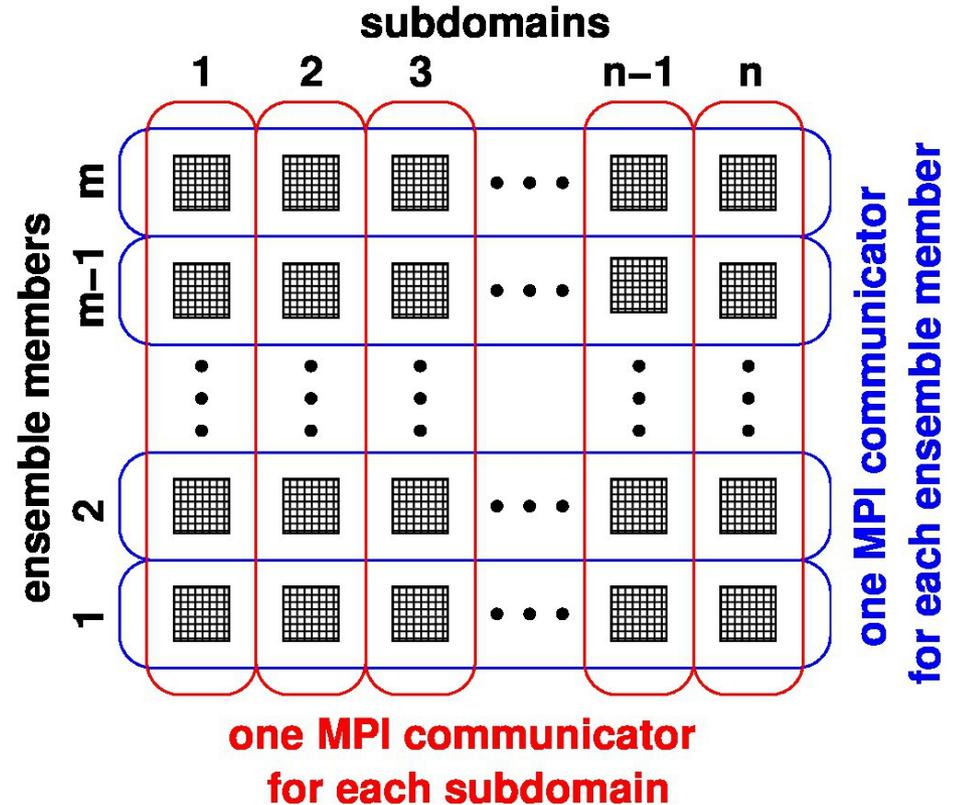
# Technological approach: ensemble simulations

Solution of the Fokker-Planck equation using a **Monte Carlo method**  
→ **ensemble NEMO simulation** → échantillon de  $p(\mathbf{x},t)$

**1 ensemble simulation**  
=  
**1 single call to NEMO**

**$m$  times more processors**  
**to run**  
 **$m$  members together in parallel**

**To each processor is given**  
**one of the  $n$  subdomains**  
**from one of the  $m$  members**



**1 MPI communicator for each member**

→ each member lives its own life, as in standard NEMO

**1 MPI communicator for each subdomain**

→ “online” computations of any feature of  $p(\mathbf{x},t)$

Uncertainty, as a key component of our systems

**What are the uncertain components  
of our systems ?**

**How to describe uncertainties ?**

**How does it participate  
to the solution of inverse problems ?**

**2**

# **Explicit simulation of uncertainties**

## 2.1 Stochastic formulation of NEMO

**Objective: transform a *deterministic* model into a *probabilistic* model**



**Describe the non-deterministic nature of the system**

**Allow objective comparison to observations**

**Introduce a weak model constraint in data assimilation**

**Method: explicitly simulate uncertainties in the mode using *random numbers***



**Propose a generic and flexible technical approach**

**Develop a first simple implementation**

**External forcing  
Unresolved scales  
Unresolved diversity**

# Autoregressive processes (1)

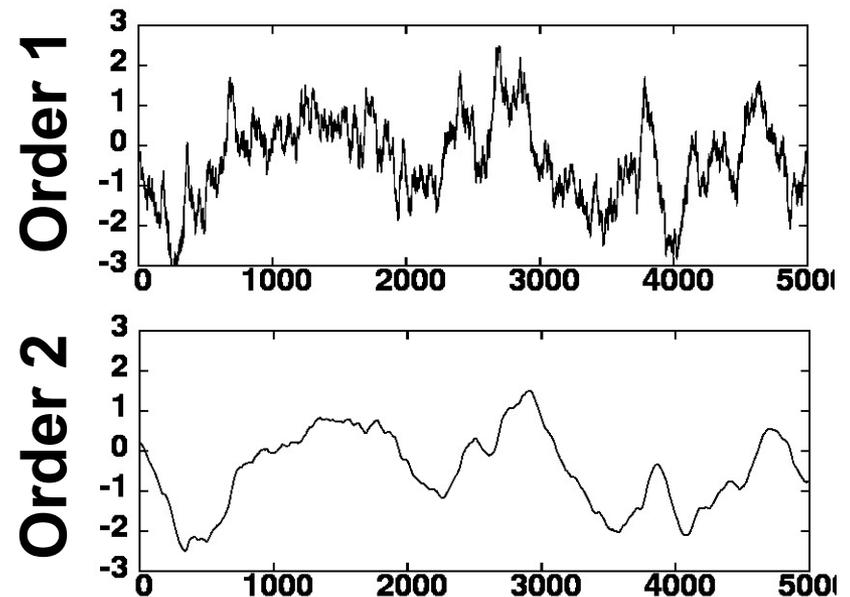
At every model grid point (in 2D or 3D), generate a set of **independent Gaussian autoregressive processes**:

$$\xi(t_k) = a \xi(t_{k-1}) + b w + c$$

where  $w$  is a Gaussian white noise ( $\rightarrow$  order 1 process)  
or an autoregressive process of order  $n-1$  ( $\rightarrow$  order  $n$  process)

Parameters  $a$ ,  $b$ ,  $c$   
to specify:

mean, standard deviation  
and correlation timescale



# Autoregressive processes (2)

## Introduce a spatial correlation structure

by applying a spatial filter to the map of autoregressive processes:

$$\tilde{\xi} = \mathcal{F}[\xi] \quad (\text{filtering operator})$$

$$\mathcal{L}[\tilde{\xi}] = \xi \quad (\text{elliptic equation})$$

which can easily be made flow dependent if needed

## Modify the marginal probability distributions

by applying anamorphosis transformation to every individual Gaussian variable:

$$\tilde{\xi} = \mathcal{T}[\xi] \quad (\text{nonlinear function})$$

for instance to transform the Gaussian variables into lognormal or gamma variables if positive noise is needed

→ This provides a generic technical way of implementing a wide range of stochastic parameterizations

# Technological approach: a new module in NEMO

These processes are generated using a **new module in NEMO**, and **can be used in any component** of the model (Brankart et al., 2015):  
circulation model, ecosystem model, sea ice model

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## Algorithm 1 sto\_par

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```
for all (map  $i = 1, \dots, m$  of autoregressive processes) do
  Save map from previous time step:  $\xi_- \leftarrow \xi_i$ 
  if (process order is equal to 1) then
    Draw new map of random numbers  $w$  from  $\mathcal{N}(0, 1)$ :
     $\xi_i \leftarrow w$ 
    Apply spatial filtering operator  $\mathcal{F}_i$  to  $\xi_i$ :  $\xi_i \leftarrow \mathcal{F}_i[\xi_i]$ 
    Apply precomputed factor  $f_i$  to keep SD equal to 1:
     $\xi_i \leftarrow f_i \times \xi_i$ 
  else
    Use previous process (one order lower) instead of white
    noise:  $\xi_i \leftarrow \xi_{i-1}$ 
  end if
  Multiply by parameter  $b_i$  and add parameter  $c_i$ :  $\xi_i \leftarrow b_i \times$ 
   $\xi_i + c_i$ 
  Update map of autoregressive processes:  $\xi_i \leftarrow a_i \times \xi_- + \xi_i$ 
end for
```

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- Generic and flexible technological approach
- Model independent implementation
- Possible to simulate many kinds of uncertainty

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## Algorithm 2 sto\_par\_init

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```
Initialize number of maps of autoregressive processes to 0:
 $m \leftarrow 0$ 
for all (stochastic parameterization  $k = 1, \dots, p$ ) do
  Set  $m_k$ , the number of maps of autoregressive processes re-
  quired for this parameterization
  Increase  $m$  by  $m_k$  times the process order  $o_k$ :  $m \leftarrow m +$ 
   $m_k \times o_k$ 
end for
for all (map  $i = 1, \dots, m$  of autoregressive processes) do
  Set order of autoregressive processes
  Set mean ( $\mu_i$ ), standard deviation ( $\sigma_i$ ) and correlation
  timescale ( $\tau_i$ ) of autoregressive processes
  Compute parameters  $a_i, b_i, c_i$  as a function of  $\mu_i, \sigma_i, \tau_i$ 
  Define filtering operator  $\mathcal{F}_i$ 
  Compute factor  $f_i$  as a function of  $\mathcal{F}_i$ 
end for
Initialize seeds for random number generator
for all (map  $i = 1, \dots, m$  of autoregressive processes) do
  Draw new map of random numbers  $w$  from  $\mathcal{N}(0, 1)$ :  $\xi_i \leftarrow$ 
   $w$ 
  Apply spatial filtering operator  $\mathcal{F}_i$  to  $\xi_i$ :  $\xi_i \leftarrow \mathcal{F}_i[\xi_i]$ 
  Apply precomputed factor  $f_i$  to keep standard deviation
  equal to 1:  $\xi_i \leftarrow f_i \times \xi_i$ 
  Initialize autoregressive processes to  $\mu + \sigma \times w$ :  $\xi_i \leftarrow \mu +$ 
   $\sigma \xi_i$ 
end for
if (restart file) then
  Read maps of autoregressive processes and seeds for the ran-
  dom number generator from restart file (thus overriding the
  initial seed)
end if
```

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## Example 1: Stochastic perturbation of parameterized tendencies

With this generic implementation, we can reproduce the SPPT scheme proposed by Buizza et al. (1999)

Separate model operator  
in NP (non-parameterized) et P (parameterized)

Assume that P is uncertain and  
simulate uncertainty by a multiplicative noise  $\xi$

$$\frac{dx}{dt} = \mathcal{NP} [\mathbf{x}, \mathbf{u}(t), \mathbf{p}] + \mathcal{P} [\mathbf{x}, \mathbf{u}(t), \mathbf{p}] \xi(t)$$

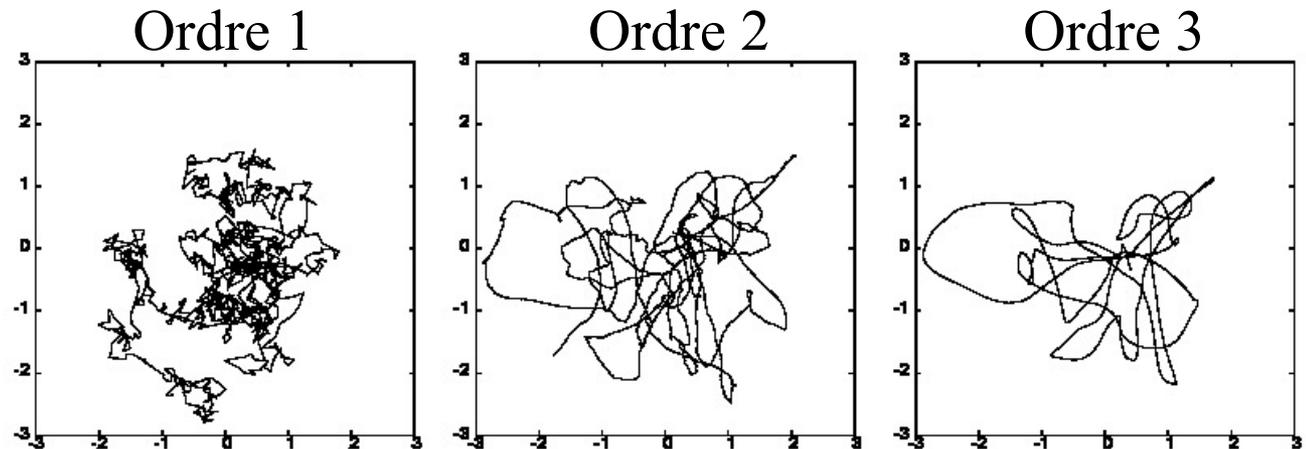
→ Use maps of autoregressive processes as  $\xi$   
(with mean 1), and specify the correlation structure  
and marginal distribution of  $\xi$ .

## Example 2: Stochastic parameterization of unresolved fluctuations

To simulate the effect of unresolved fluctuations in the nonlinear terms of the model equations

$$\frac{d\mathbf{x}}{dt} = \frac{1}{n} \sum_{i=1}^n \mathcal{M} [\mathbf{x} + \delta\mathbf{x}_i(t), \mathbf{u}(t), \mathbf{p}] \quad \text{with} \quad \sum_{i=1}^n \delta\mathbf{x}_i(t) = 0$$

Generate fluctuations using **random walks** around every grid point



→ Use maps of autoregressive processes as components  $x, y, z$  of the random walks. Specify space and time correlation structure.

## Example 3: Stochastic parameterization of unresolved diversity

**To simulate the effect of the unresolved diversity of system behaviours (e.g. biological diversity,...)**

This assumes that the system simultaneously includes a variety of possible behaviours, which cannot be described by one single value of each parameter.

$$\frac{d\mathbf{x}}{dt} = \frac{1}{n} \sum_{i=1}^n \mathcal{M}[\mathbf{x}, \mathbf{u}(t), \mathbf{p} + \delta\mathbf{p}_i(t)]$$

For instance, the ecosystem usually includes many different species of phytoplankton and zooplankton, each with its own behaviour, while the model can only resolve a few classes of species.

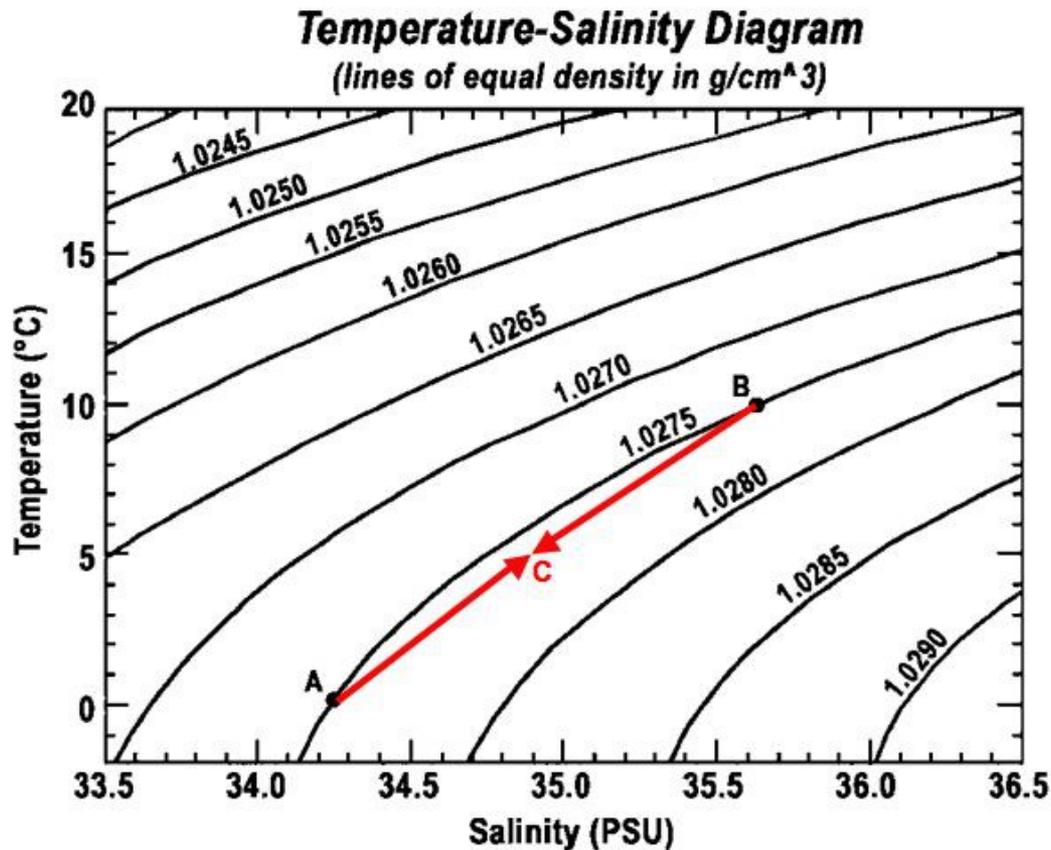
**→ Use maps of autoregressive processes  $\xi$  as multiplicative noise for the parameters, and specify their correlation structure and their marginal distribution.**

**3**

# **Stochastic circulation model**

# Uncertainties in the computation of density

In the model, the large-scale density is computed from large-scale temperature and salinity, using the sea-water equation of state.



(a)

Mixing waters of equal density but different T&S systematically increases density (cabbeling)

(b)

Averaging T&S equations systematically overestimates density (in a fluctuating, non-deterministic way)

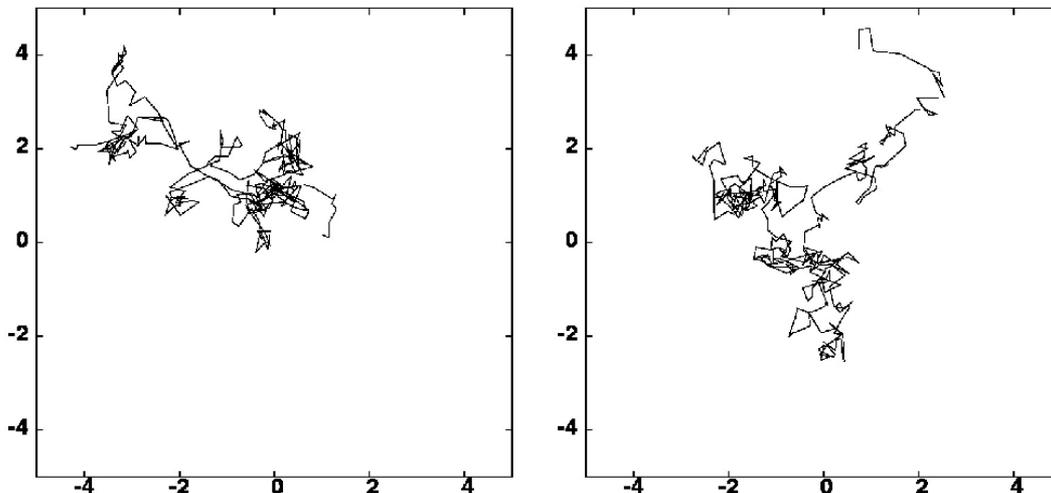
Because of the nonlinearity of the equation of state, unresolved scales produce an average effect on density.

# Random walks to simulate unresolved temperature and salinity fluctuations

Computation of the random fluctuations  $\Delta T_i$  et  $\Delta S_i$   
as a scalar product of the local gradient with random  
walks  $\xi_i$

$$\Delta T_i = \xi_i \cdot \nabla T \quad \text{and} \quad \Delta S_i = \xi_i \cdot \nabla S$$

## Random walks



## Assumptions

AR1 random processes

uncorrelated on the horizontal

fully correlated  
along the vertical

5-day time correlation

horizontal std: 2-3 grid points  
vertical std: <1 grid point

# Stochastic equation of state for the large scales

## Stochastic parameterization (Brankart, 2015)

using a set of random T&S fluctuations

$$\Delta T_i \text{ et } \Delta S_i, i=1, \dots, p$$

to simulate unresolved T&S fluctuations

$$\rho = \frac{1}{p} \sum_{i=1}^p \rho [T + \Delta T_i, S + \Delta S_i, p_0(z)] \quad \text{with} \quad \sum_{i=1}^p \delta T^{(i)} = 0, \quad \sum_{i=1}^p \delta S^{(i)} = 0$$

No effect if the equation of state is linear.

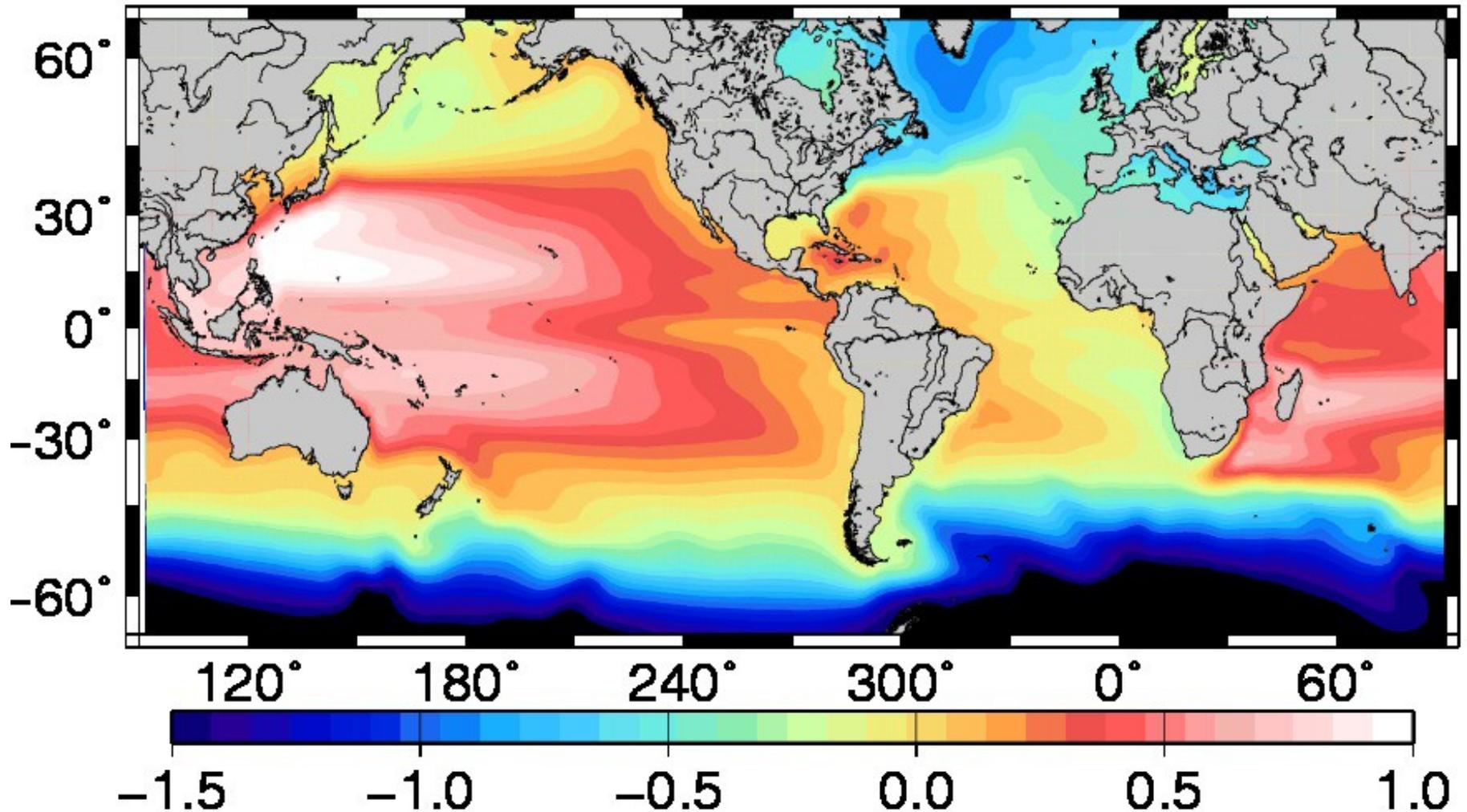
Proportional to the square of unresolved fluctuations.

## Correction $\Delta\rho$ applied in the thermal wind equation, as in the semi-prognostic method of Greatbatch et al. (2004)

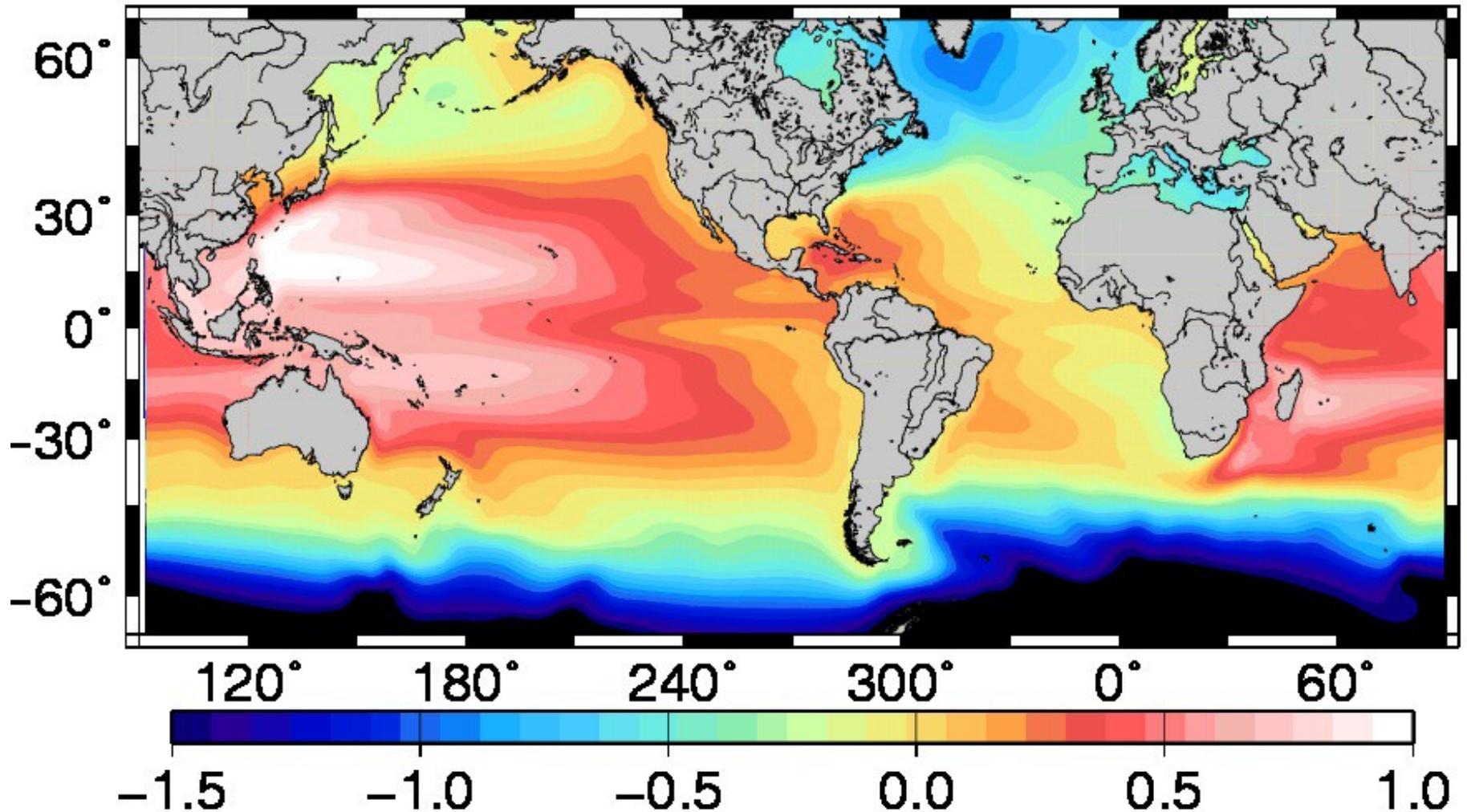
No direct modification of T&S; no enhanced diapycnal mixing.

T&S only modified indirectly through a modification  
of the main currents

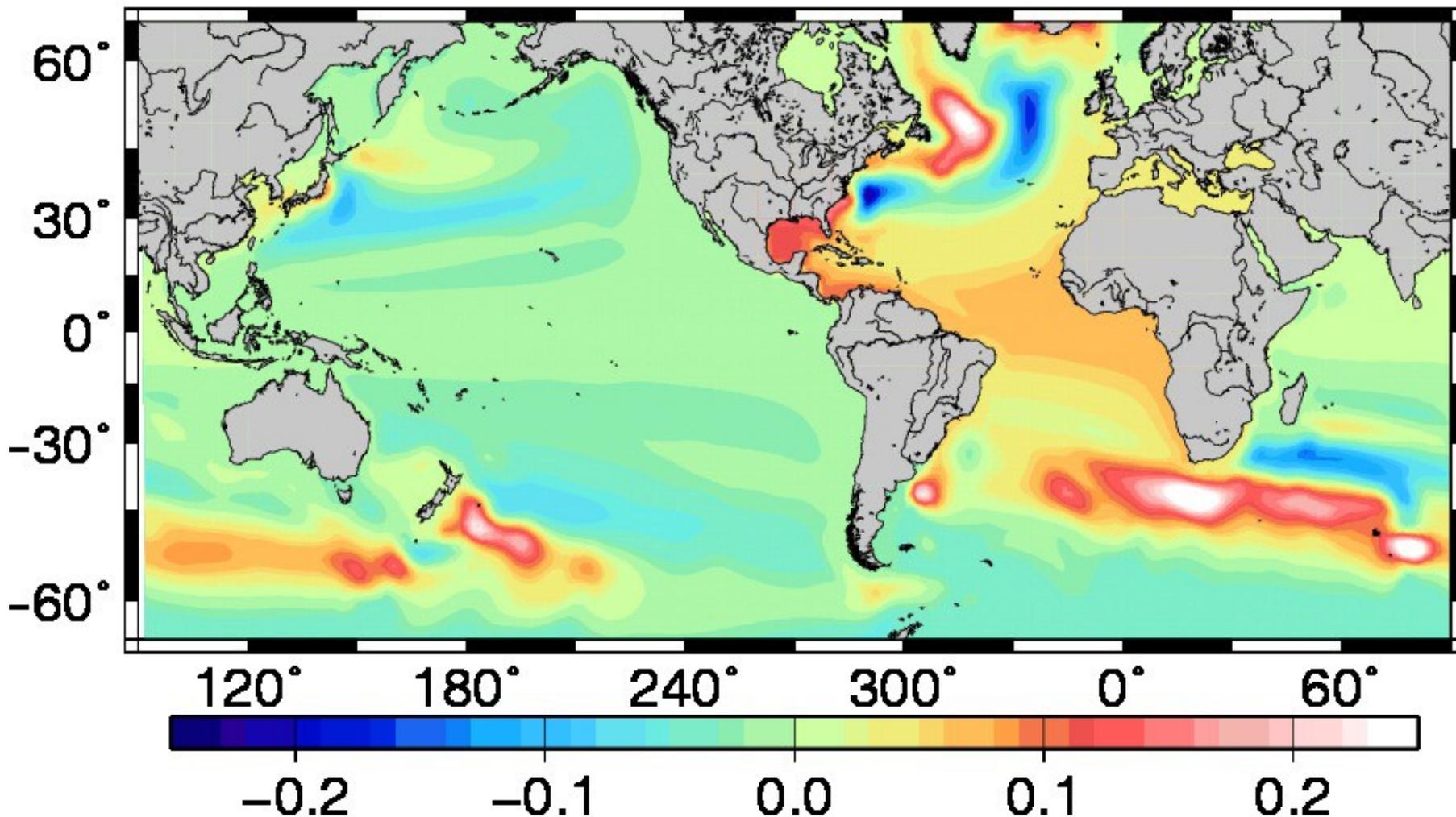
# Mean sea surface elevation (standard)



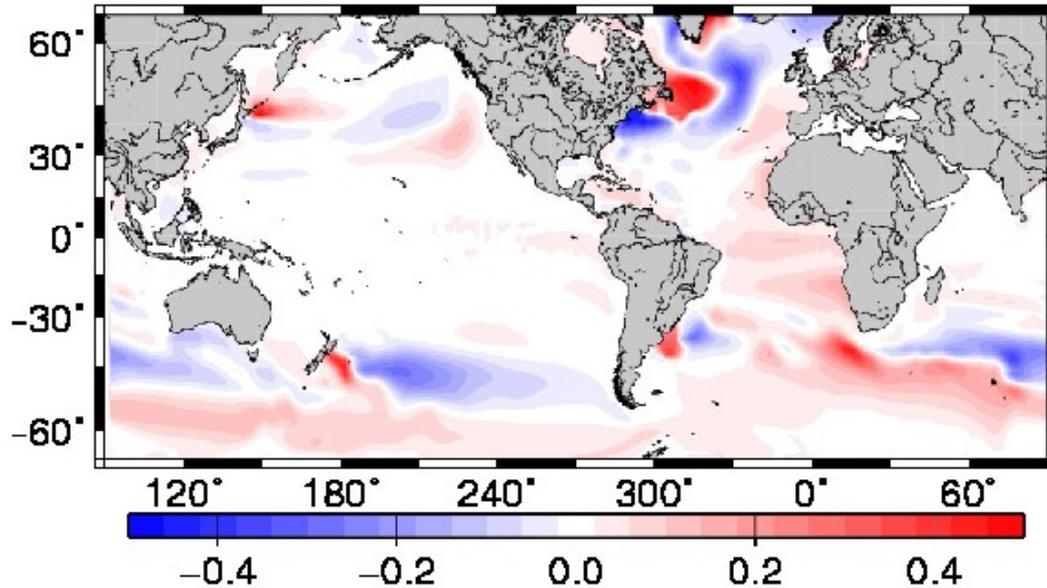
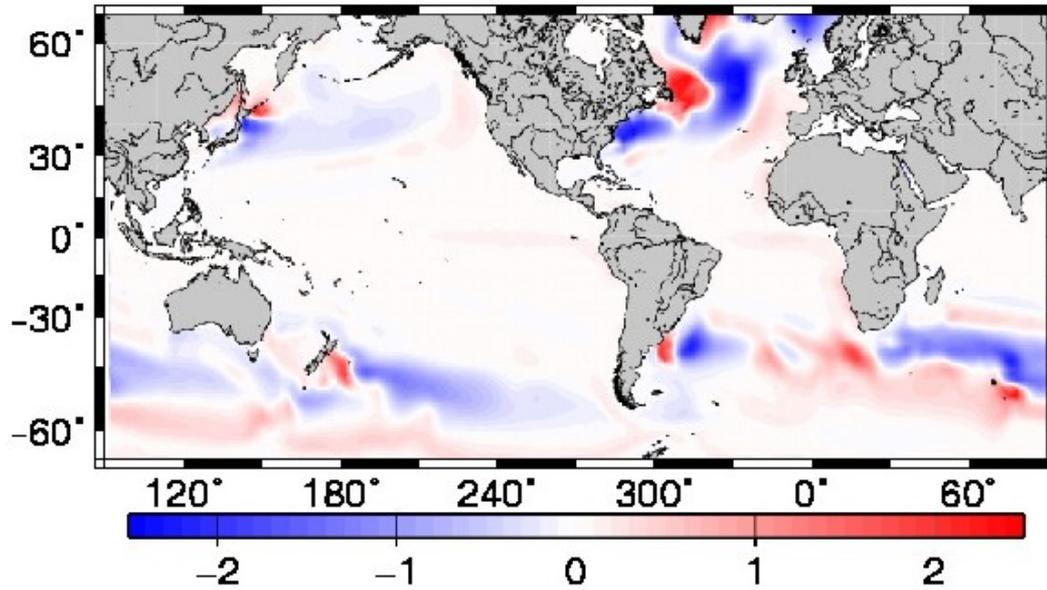
# Mean sea surface elevation (stochastic)



# Mean sea surface elevation difference



# Averaged SST & SSS difference



**Modification  
of the mean flow**



**Modification  
of the mean  
SST & SSS**

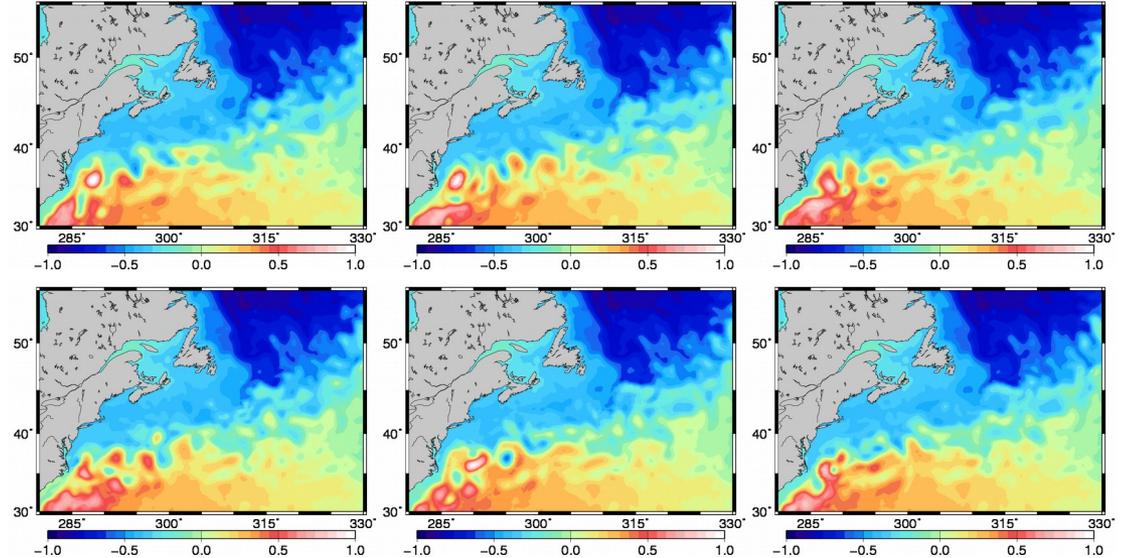


**Modification  
of air/sea  
interactions**

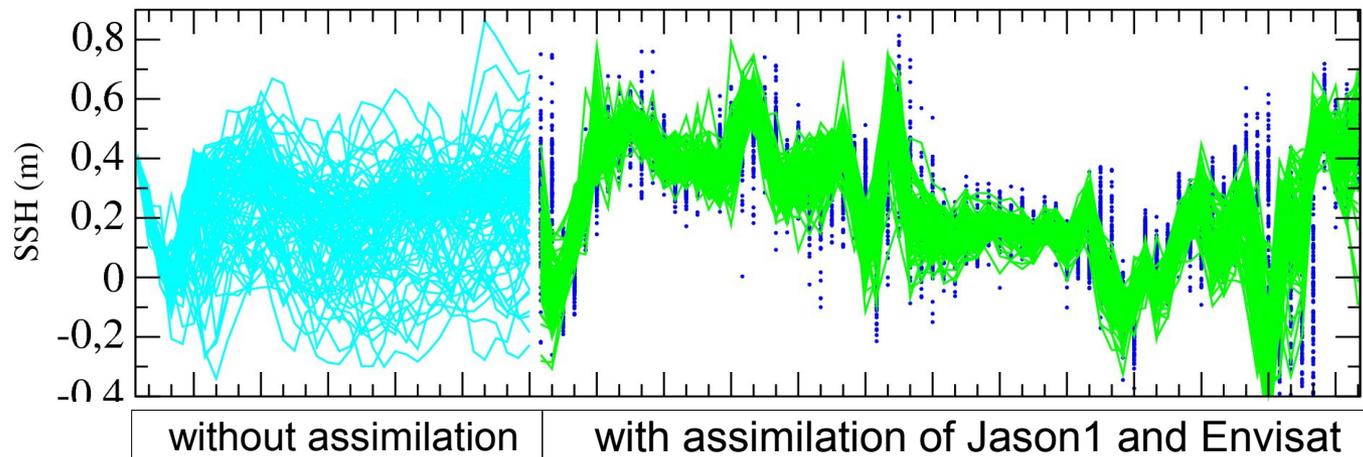
# Ensemble of mesoscale flows

## Probability distribution for SSH

as simulate here by ensemble NATL025 (large case SANGOMA benchmark): 6 members among 96



→ **assimilation of altimetric data (in SANGOMA, Candille et al., 2014)**



Time evolution of the pdf

from June 2005 to December 2006

**4**

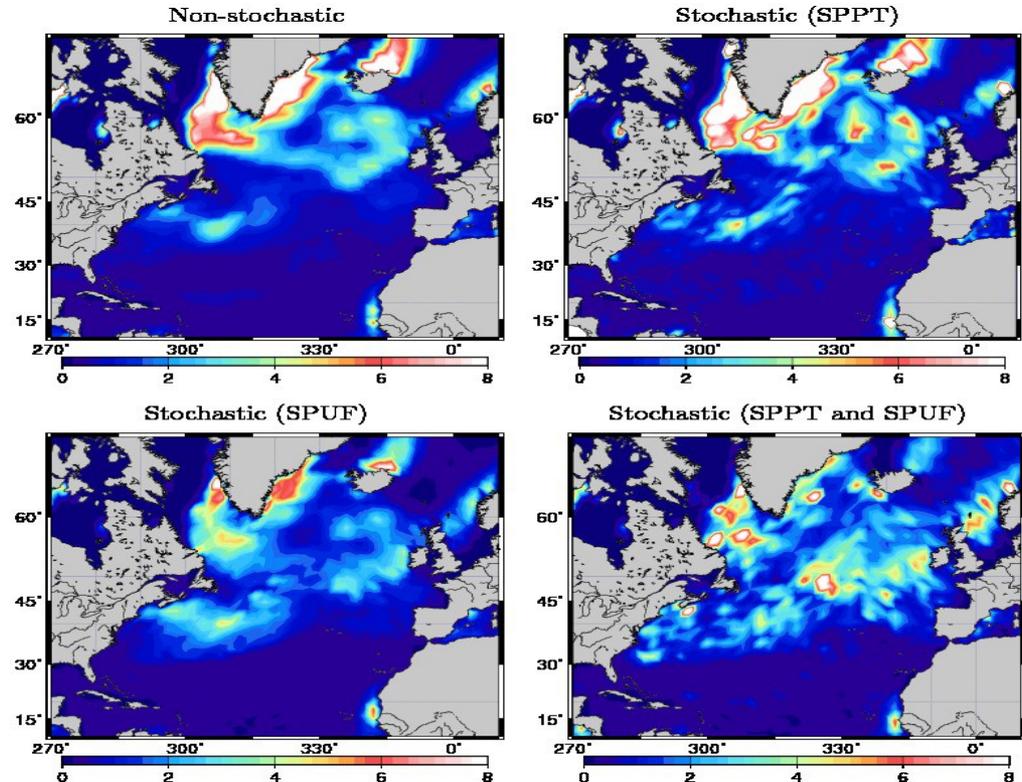
# **Stochastic ecosystem model**

# Stochastic ecosystem model

**Multiple sources of uncertainty in ecosystem model:  
unresolved biological diversity, unresolved scales, etc.**

**Unresolved diversity  
multiplicative noise  
in the SMS terms  
of the model**

**Unresolved scales  
stochastic processes  
explicitly simulating  
unresolved  
fluctuations of  $C_i$**



- Considerable effect on the mean behaviour of the system
- Increase of patchiness (↔ ocean colour data)

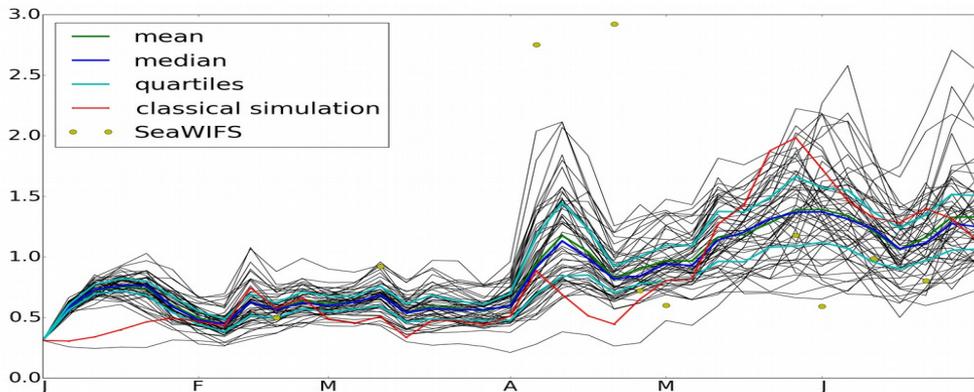
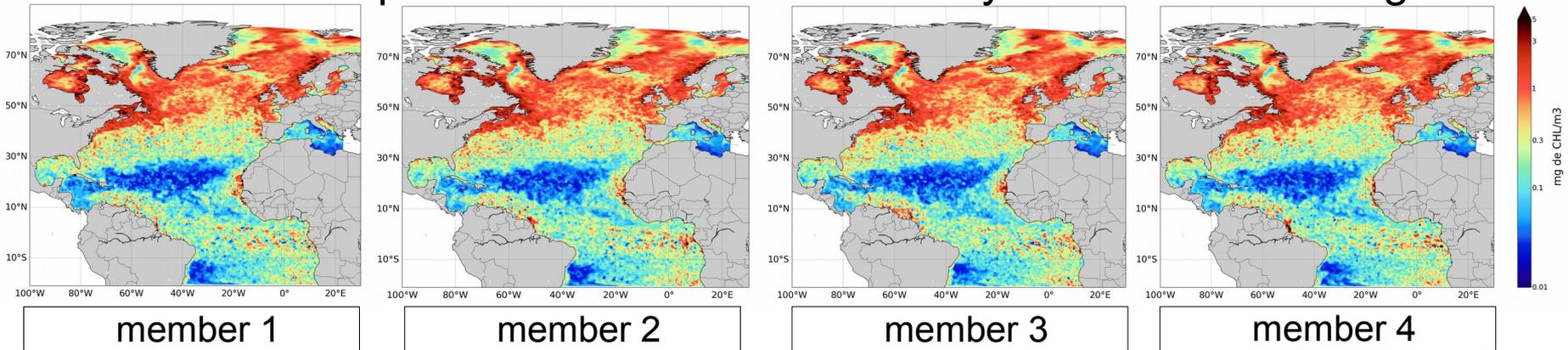
# Ensemble simulation of the ecosystem

## Probability distribution of chlorophyll concentration

as simulated here by ensemble NATL025/PISCES

(large case Sangoma benchmark)

with stochastic parameterization of uncertainty: 4 members among 50



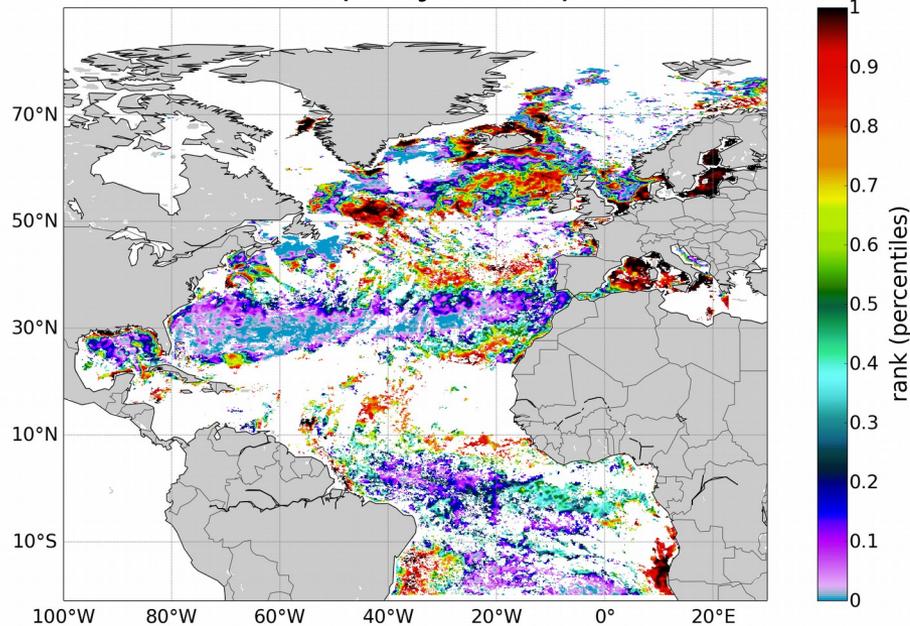
Time evolution  
of the pdf  
for phytoplankton

from January to June 2005

→ assimilation of ocean colour observations  
(projects FP7-MyOcean2 and SANGOMA, Garnier et al., 2015)

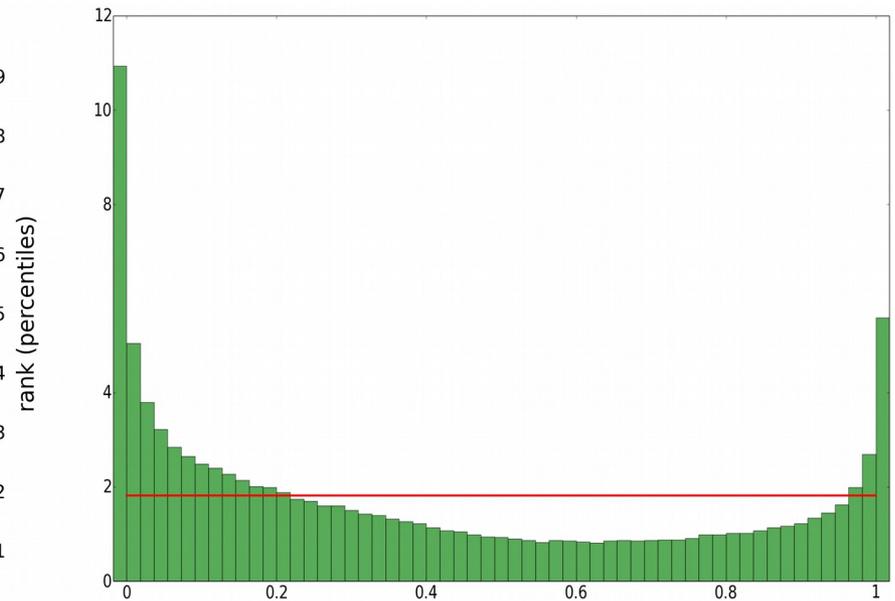
# Comparison to ocean colour observations

## Rank of SeaWifs observations in the ensemble simulation (May 2005)



The ensemble spread is already sufficient to include more than 80% of the observations (accounting for a 30% observation error)

## Rank histogram for SeaWifs over the whole domain



The ensemble is not far from being reliable, even if still underdispersive (too many observations in the external ranges of the ensemble)

- objectively test consistency between simulations and observations
- prerequisite to ocean colour **data assimilation de données**

**5**

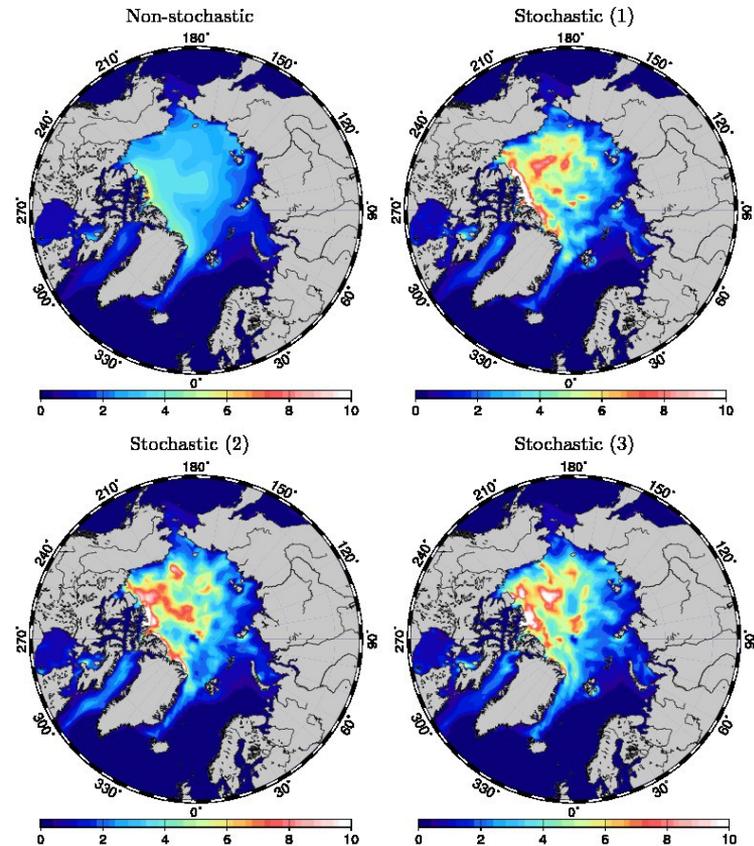
**Stochastic sea ice model**

# Stochastic sea ice model

An important difficulty of sea ice model  
is the unresolved diversity of dynamical sea ice behaviours

One of the most sensitive  
sea ice parameters  
is ice strength ( $P^*$ ):

multiplicative noise  
applied to  $P^*$   
(parametrization of  
Juricke et al., 2013,  
implemented in NEMO)



- Considerable effect on the mean ice thickness (in ORCA2)
- Intrinsic interannual variability is stimulated

**6**

# **Conclusions**

**The NEMO model becomes probabilistic;  
it is seen as a complex system,  
built up from uncertain components**

→ The goal of ocean modelers is then to build a model  
as informative as possible at the lesser cost.

**This probabilistic description requires  
ensemble simulations**

- Objective comparison between simulations and observations
- Deal with model uncertainty in ocean data assimilation systems

An appropriate simulation of uncertainty is necessary to make the link between model, observations, and data assimilation systems

**Uncertainty is bound to become a key constituent of the systems that we are using in oceanography, not something that can be thought separately from the results**

**Properly dealing with uncertainty will require an integrated engineering approach at the interface between oceanography and applied mathematics**